

Sutures, Dyke Swarms and Carbonatites in People's Democratic Republic of Yemen

M.A. BA-BTTAT^{*}, B.F. WINDLEY^{**},

A.T. AL-MISHWIT^{*}, and D.C. ALMOND^{*}

^{*}*Geology Department, Kuwait University, Kuwait, and* ^{**}*Geology
Department, Leicester University, Leicester, U.K.*

ABSTRACT. The basement rocks in the Lowder-Mudiah area of the Abyan Governorate, P.D.R.Y. (South Yemen) consist of three major lithological units. These are, *viz.*, granitic gneisses, metavolcanics and dyke swarms, with their granite diorite and gabbro host rocks, unconformably and partly covered by Jurassic limestones and Quaternary basalts.

These three lithological units are separated by two main, northeast striking, ductile shear zones, along which sutures are believed to have taken place. Carbonatite dykes intruded through gray granitic gneiss host rock.

Introduction

The Lowder-Mudiah area is situated within the Abyan Governorate of the People's Democratic Republic of Yemen (P.D.R.Y.) (Fig. 1), between longitudes 45°40' and 46°10'E and latitudes 13°40' and 14°00'N. The area covers approximately 2,200 sq. km and is underlain by a basement of granitic gneiss and a sequence of metavolcanic rocks, regarded by Greenwood and Bleackley (1967) as a part of the Aden Metamorphic Group. These rocks are unconformably overlain by Jurassic limestone (Greenwood and Bleackley 1967) and Quaternary volcanic basalt of the Shugra volcanic field (Cox *et al.* 1977). Elevation in the mapped area varies between 1,100 and 1,200 m, except at the Mukeras Escarpment, where elevations in the mapped area exceed 2,500 m. The escarpment separates the Lowder-Mudiah lowlands from the Mukeras Plateau to the north (Fig. 2).

This paper is part of an ongoing Ph.D. research project by the first author. Mapping is based on aerial photographs at a scale of 1:60,000.

General Geology

The geology of the area is dominated by Precambrian basement rocks, which are unconformably overlain by Jurassic limestone and partly covered by Quaternary basalt flows of the Shugra volcanic field. The basement rocks consist of three major lithological units that are separated by two northeast-striking,

ductile-shear thrust zones. The three lithological units are: a) gneissose granites underlying a central belt, b) a sequence of bimodal metavolcanic rocks in the southeast, and c) a weakly altered white granite, diorite and gabbro intruded by mafic to felsic dyke swarms to the northeast. A ductile shear (thrust) zone is located within quartz-biotite gneiss and amphibolites of the central belt. Localized migmatite zones were also found within the same quartz-biotite gneiss and amphibolites.

Sulphide gossans are developed on mineralized quartz veins in the metamorphosed quartz-rhyolite porphyry. Chemical analysis of surface samples reveal that they contain traces of Ag, Pb and Zn.

Pb-Zn mineralization in the form of cavity filling (Karst), associated with barite, is hosted by the Jurassic limestone. It is found in certain pockets outside the border of the mapped area and seems to resemble the Pb-Zn (Mississippi-type) mineralization in the Jurassic (Amran) limestone in Nehim (Jibelah) of North Yemen (Al-Shatory 1989, personal communication).

Structural Geology

Multiple deformation has affected the area and is represented by three recognisable fold episodes, each with its own distinctive style and orientation. North-east trending isoclinal folds were followed by open folds with vertical axial planes and by northwest trending cross- and kink-folds.

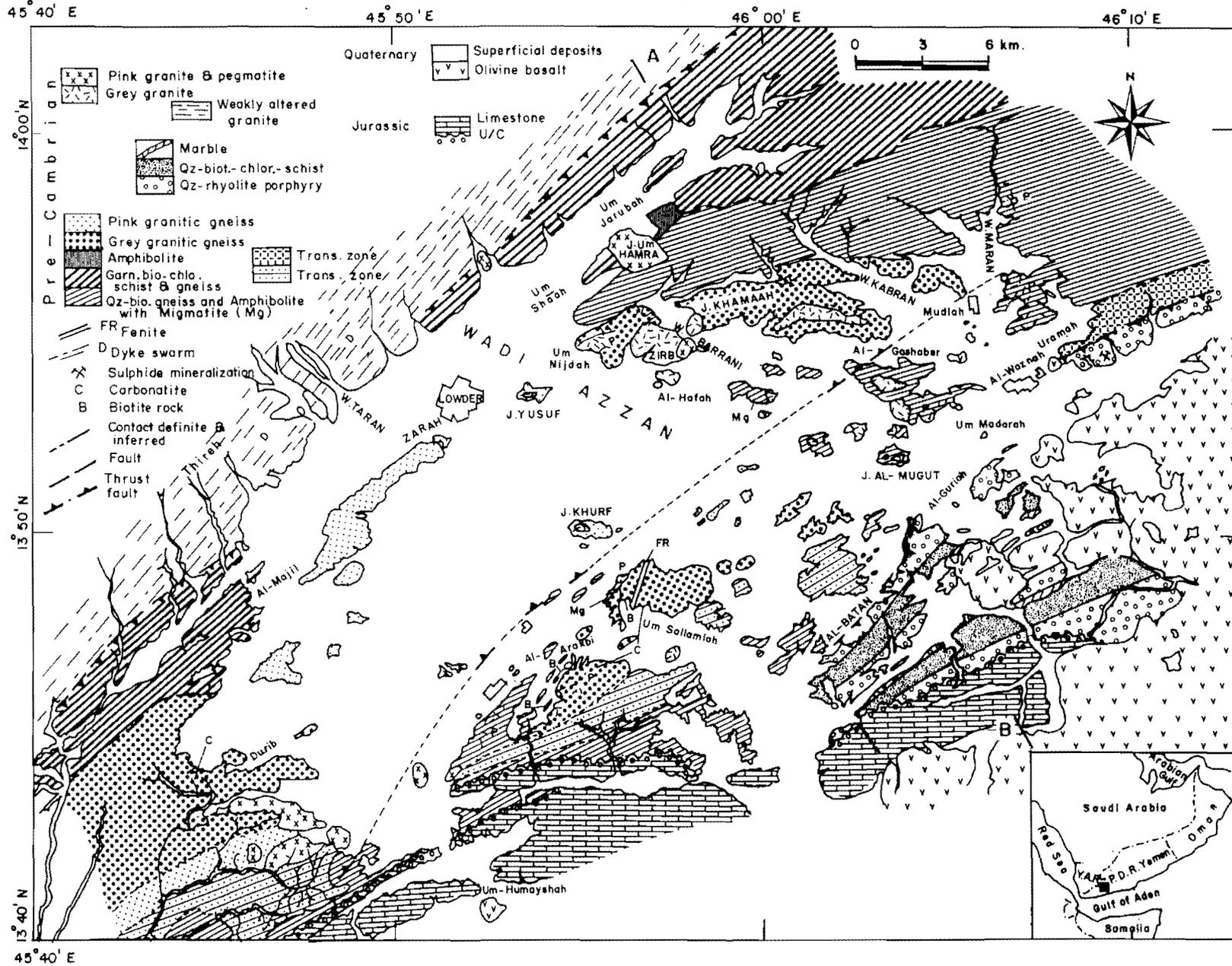


FIG. 1. Geological map of Lower-Mudiah area (P.D.R. Yemen).

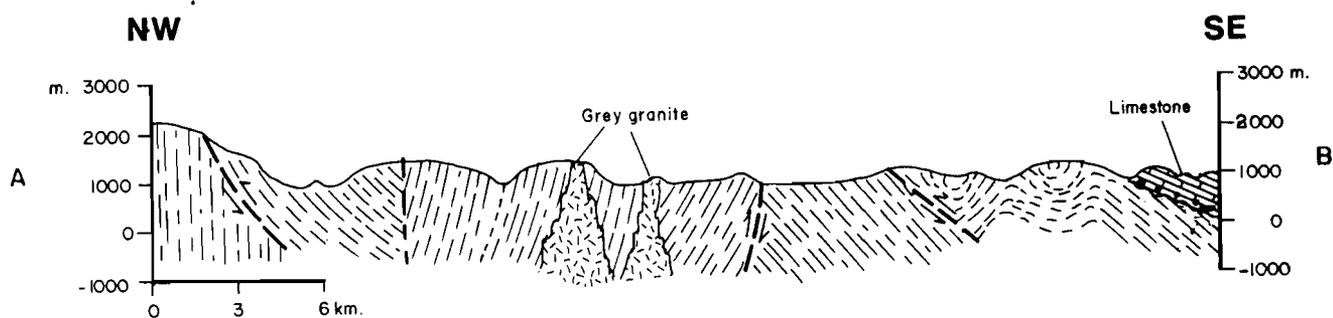


FIG. 2. Structural cross-section along A.B.

Two main ductile shear zones developed on major thrusts, striking northeast. The first zone is located in the eastern part of the area and separates the metavolcanic lithological unit to the east from the granitic gneisses and amphibolites in the central part of the area. The second zone is located in the western margin of the area, separating granitic gneisses and amphibolites from the dyke swarms and their granite, diorite and gabbro host rocks. Shearing along these zones is evidenced by mylonitization, lensoidal fragmentation and sigmoidal textures as characteristic shear fabrics.

A third major shear (thrust?) zone is located within the gneiss belt, marked by sheared quartz-biotite gneiss and amphibolite which have been folded into a northeast plunging antiform. Dislocation is indicated by strong shearing in the quartz-biotite gneisses and amphibolite and their disruption into lens-shaped fragments, and by the interlayering of slices of different rock types. The above assumptions about thrust follow the description by Park (1983).

Sutures

Shackleton (1986) suggested that the dismemberment of ophiolites into lenses and repetition of its units imply the presence of major crustal shear zones. If this suggestion is valid, then the presence of lenses of dismembered, tectonized ophiolite about 100 km southwest of the Lowder-Mudiah area (Al-Derweesh 1988) and along the strike of the ductile shear zone may indicate the presence of a NE-trending suture zone. Another basic-ultrabasic and serpentinitized sequence at Mukeras, to the Northwest of Lowder-Mudiah (Dobrenky *et al.* 1977, Saeidan 1989, personal communication) may be indicative of a parallel suture, since it is situated along a major lineament passing through both South and North Yemen.

The proposed sutures seem to be more acceptable if one considers the recent findings of the presence of a suture zone in the horn of Africa and Ethiopia (Windley 1988, personal communication), since the

latter region can be regarded as a continuation of South Arabia. Furthermore, similar suture zones parallel to those in South Arabia are under investigation in Madagascar (Windley 1988, personal communication).

Dyke Swarms

The third lithological unit of the basement occupies the western part of the area and forms a northeast trending, 1 km high escarpment that reaches elevations of 2,500 m. The main rock unit is the dyke swarm, which underlies an area of about 20 × 60 km. The dykes are bimodal in composition and intrude rocks which vary from weakly altered white granite to diorite (740 ± 22 Ma) and gabbro. The dikes occupy at least 76% of the total rock volume with an average thickness of 10 m each, and are Precambrian in age since they give K-Ar apparent age of (587 ± 18 - 709 ± 21 Ma). Trace element analysis show that they have moderate Nb content, and in this and other respects they are similar to intra-continental flood basalts, as also are the basaltic sills within the metavolcanic lithological unit. These rocks may indicate an extensional tectonic environment. The host granitic rock is calc-alkaline, as are the granitic gneisses and the late gray, pink and pegmatite granites; such rocks are likely to be subduction related. Windley and Tarney (1986), discussing the structural evolution of the lower crust in general, stated that exposed basic dyke swarms often represent feeders of lavas within aulacogens under continental forelands. The Mukeras Escarpment dykes may mark a major late Precambrian rifting episode in this segment of the Arabian Shield.

Carbonatite

Carbonatite occurs as dykes intruded into the grey granitic gneiss in the vicinity of Um-Sallamiah, Al-Arakbi and Durib. The dykes are associated with uranium, thorium and potassium anomalies (Abouy *et al.* 1981). Trace and Rare Earth Element analysis support and confirm field evidence that they are carbonatite, rather than marble.

Conclusion

Although some features of the geology of Lowder-Mudiah are similar to those of the Precambrian of the Arabian Shield in Saudi Arabia (Greenwood *et al.* 1980, Robool *et al.* 1983, and Stoesser *et al.* 1983), there are closer comparisons, in both lithology and structure, with the Precambrian basement of north Somalia (Ba-bttat *et al.* 1990). Thus, there is a possibility that the Yemeni basement is distinct from that of Saudi Arabia and represents a northeastern continuation of the Mozambique Belt, which swings round from north-south to northeast-southwest into southwest Arabia (Warden and Daniels 1983, and Warden and Horkel 1984). A small segment of the Pan-African Shield in Dhofar (Hawkins *et al.* 1981) may also be a part of this belt. The Lowder-Mudiah rocks could represent early Pan-African deep crust thrust up to shallow levels in an island-arc environment. The volcanics of Saudi Arabia (Greenwood *et al.* 1980, Camp 1984, Al-Shanti and Mitchell 1976, Ba-bttat 1978, 1981 and 1985, and Ba-bttat and Hussein 1983) may represent the arc-rock themselves. In other words, the dyke swarms and their granitic host could lie along the southern edge of the Saudi Arabian island-arc belt, while the granitic gneiss of Lowder-Mudiah were thrust up between the Mozambique and Saudi Arabian Terrains.

Acknowledgement

The authors are very grateful to the government of the People's Democratic Republic of Yemen, Ministry of Oil and Minerals, for helping field works, to the British Council for granting a Foreign and Commonwealth Office Award Scholarship, and also Kuwait and Leicester Universities. Thanks are due to David C. Rex of the University of Leeds for carrying out the K-Ar analysis for age determination of the dyke swarms and their country rocks. Thanks are also due to the technical staff members of the Departments of Geology at both Kuwait and Leicester Universities.

References

- Abouv, A.K., Lysitsin, E.S., Potroukhin, I.I., Lychkovsky, B.P. and Abdul Aziz, I.S. (1981) *Report on the results of prospecting and estimation within the Lawder field of carbonatites for 1977-1981*, Technical Report, Board of Petroleum and Minerals, Dept. of Geol. and Miner. Exploration, Aden, P.D.R.Y. (unpublished).
- Al-Derweesh, K.A.S. (1988) *Geology and Hydrogeology of Batais Al-Husn area (Upper Abyan Delta) Abyan Governorate, P.D.R. of Yemen*, M.Sc. Thesis, Kuwait University, Kuwait (unpublished).
- Al-Shanti, A.M. and Mitchell, A.H.G. (1976) Late Precambrian subduction and collision in the Amar-Idsas region, Arabian Shield, Kingdom of Saudi Arabia. *Tectono physics* **30**: T41-T47.
- Ba-bttat, M.A. (1978) *The Geology of Wadi Al-Maragh area, Saudi Arabia*. High Diploma Thesis, Fac. Earth Sci., King Abdulaziz Univ., Jeddah, Saudi Arabia (unpublished).
- . (1981) *Geology and Mineralization of the Jabal Samran and Jabal Abu-Mushut area*, M.Sc. Thesis, Fac. Earth Sci., King Abdulaziz Univ., Saudi Arabia (unpublished).
- . (1985) The Geology of Wadi Al-Maragh area, Saudi Arabia. In: *The Proceedings of the 2nd Jordanian Geological Conference, Amman, Jordan*, pp. 428-456.
- Ba-bttat, M.A. and Hussein, A.A. (1983) Geology and Mineralization of the Jabal Samran and Jabal Abu-Mushut area. *Bull. Fac. Earth Sci.*, King Abdulaziz Univ., **6**: 571-578.
- Ba-bttat, M.A., Windley, B.F., Al-Mishwit, A.T. and Almond, D.C. (1990) Geology of Lowder-Mudiah area, Abyan Governorate, People's Democratic Republic of Yemen. In: *The Proceedings of the International Conference on the Geology of Somalia and the Surrounding Regions (GEOSOME 87), Special issue of the Journal of Africa and the Middle East* (in press).
- Camp, V.E. (1984) Island arcs and their role in the evolution of the Western Arabian Shield. *Geol. Soc. Am. Bull.* **95**: 913-921.
- Cox, K.G., Gass, I.G. and Mallick, D.I.J. (1977) The western part of the Shugra volcanic field, South Yemen. *Lithos* **10**: 185-191.
- Dobrenky, A.E., Arutunyan, G.S., Bolkovoy, B.I., Agaronyan, L.V. and Pogosyan, G.A. (1977) *Geological Report on Results of Search and Estimation at Mukeyras deposit*, B.P.M., Dept. of Min. and Explor. P.D.R. of Yemen (Aden) (unpublished).
- Greenwood, J.E.G.W. and Bleackley (1967) Geology of the Arabian Peninsula. Aden Protectorate, *U.S.G.S. Professional Paper*, **560-C**.
- Greenwood, W.R., Anderson, R.E., Fleck, R.J. and Roberts, R.J. (1980) Precambrian Geological History and Plate Tectonic Evolution of the Arabian Shield, Jeddah, Saudi Arabia. *D.G.M.R. Bull. No. 24*.
- Hawkins, T.R.W., Hindle, D. and Strugnell, R. (1981) Outlines of the Stratigraphy and Structural Framework of South Dhofar, Sultanate of Oman. *Geologieen Mijinun* **60**: 247-256.
- Park, R.G. (1983) *Foundations of Structural Geology*, Blackie Son Ltd., Glasgow, Scotland, U.K.
- Robool, M.J., Ramsay, C.R., Jackson, N.J. and Darbyshire, D.P.F. (1983) Petrological aids to stratigraphic correlation of volcano sedimentary successions in the central Arabian Shield. *Bull. Fac. Earth Sci., King Abdulaziz Univ.* **6**: 165-194.
- Shackleton, R.M. (1986) Precambrian collision tectonics in Africa. In: Coward, M.P. and Ries, A.C. (ed.), *Collision Tectonics*, Blackwell, Scientific Oxford, U.K., pp. 329-349.
- Stoesser, D.B., Fleck, R.J. and Stacey, J.S. (1983) Geochronology and origin of an early tonalite gneiss of the Wadi Tarib batholith and the formation of syntectonic gneiss complexes in the southeastern Arabian Shield, Kingdom of Saudi Arabia. *Bull. Fac. Earth Sci., King Abdulaziz Univ.* **6**: 351-364.
- Warden, A.J. and Daniels, J.L. (1983) Evolution of the Precambrian of Northern Somalia. *Bull. Fac. Earth Sci., King Abdulaziz Univ.* **6**: 145-164.
- Warden, A.J. and Horkel, A.D. (1984) The Geological Evolution of the NE-Branch of the Mozambique Belt (Kenya, Somalia, Ethiopia). *Mitt. Osterr. Geol. Ges.* **78**.
- Windley, B.F. and Tarney, J. (1986) The structural evolution of the lower crust of orogenic belts, present and past. In: Dawson, J.B., Carswell, D.A., Hall, J. and Wedepohl, K.H. (eds), *The Nature of the Lower Continental Crust*, Geological Society, London, Special Publication No. 24, pp. 221-230.

دروز التحام وحشود جُدديّة قاطعة وصخور كربوناتيتية في جمهورية اليمن الديمقراطية الشعبية

محفوظ بابطاط* ، ب . ف . وندي** ، علي المشوط* و د . سي . ألموند*
* قسم الجيولوجيا ، جامعة الكويت ، الكويت ؛ و ** قسم الجيولوجيا ، جامعة ليستر ،
المملكة المتحدة

المستخلص : تشتمل منطقة لودر - مودية من محافظة أبين بالجمهورية اليمنية على ثلاث وحدات صخرية رئيسية ، هي : نيس جرانيتي وبركانيات متحولة وحشد من الجُدّد القاطعة مع صخر مُضيف من الديوريت الجرانيتي والجابرو . يُغطي هذه الوحدات الصخرية جزئياً ولا توافقاً حجر جيرى جوراسي وبازلت ثلاثي .
يفصل هذه الوحدات الصخرية عن بعضها البعض نطاقاً جَزْراً ، ينحوان باتجاه الشمال الشرقي ، ويُعتقد بأنه قد حدث على طولها درز التحام .

كما تحللت صخور النيس الجرانيتي الرمادية المضيفة جُدّد قاطعة من صخور الكربوناتيت .